

Improving Short-Term QPFs over Complex Terrain with a WRF-LETKF Radar Data Assimilation System: OSSEs on Typhoon Morakot (2009) Chih-Chien Tsai, Shu-Chih Yang, and Yu-Chieng Liou Department of Atmospheric Sciences, National Central University, Taiwan

### **1. Introduction**

This study implements a Doppler radar observation operator for the WRF-LETKF data assimilation system (Yang et al. 2012). OSSEs on Typhoon Morakot (2009) are performed to investigate the multivariate interactions in this system with a goal to optimize the assimilation strategies for improving the short-term QPF over complex terrain.

# 2. Methodology

LETKF: Local Ensemble Transform Kalman Filter (Hunt et al. 2007)
 1) Simultaneously update the ensemble mean (state) and perturbations

## 4. Results and discussions

- Single-point assimilation: P and Q (Fig. 4) are located to the due west and due north of the radar. They contain the most complete u and v
  - information, respectively.
  - 1) Corresponding fields are broadly corrected
  - 2) Patterns of *w* and  $q_r$  are
  - noisier than *u* and *v*
  - $\rightarrow$  scale difference
  - 3) P is better than Q for its higher signal-to-noise



(uncertainty) locally in space by  $\bar{\mathbf{x}}_{a} = \bar{\mathbf{x}}_{f} + \mathbf{X}_{f} \tilde{\mathbf{P}}_{a} \mathbf{Y}_{f}^{T} \mathbf{R}^{-1} (\mathbf{y}_{o} - \bar{\mathbf{y}}_{f})$   $\mathbf{X}_{a} = \mathbf{X}_{f} [(K - 1) \tilde{\mathbf{P}}_{a}]^{1/2}$   $\tilde{\mathbf{P}}_{a} = [(K - 1)\mathbf{I}/\rho + \mathbf{Y}_{f}^{T} \mathbf{R}^{-1} \mathbf{Y}_{f}]^{-1}$ 

- $\overline{\mathbf{x}}$ ,  $\mathbf{X}$ : ensemble mean and perturbations  $\overline{\mathbf{y}}$ ,  $\mathbf{Y}$ : ensemble mean and perturbations in the observation space
  - **R** : observation error covariance matrix
- $\mathbf{y}_o$ : observations K: ensemble size
- $\rho$  : covariance inflation factor

2) Grid points can be processed in parallel for computational efficiency.

#### Model setup:

- 1) Model: WRF-ARW V3.2.1
- 2) Physics:
  - a) Purdue Lin microphysics
  - b) Kain-Fritsch cumulus parameterization
  - c) Noah land-surface model
  - d) YSU planetary boundary layer
- **3)** <u>IC and BC</u>: NCEP  $1^{\circ} \times 1^{\circ}$  FNL

#### Radar observation operator:

#### 1) Spatial conversion:

- a) Vertical interpolation to the intersections
  - of the sweeps and grid columns
- b) Consider the earth surface curvature, atmospheric refraction and terrain
- 2) Variable conversion:
  - a) Radial velocity:  $V_r = [ux + vy + (w v_t)z](x^2 + y^2 + z^2)^{1/2}$ where  $v_t = 5.40(p_0/\bar{p})^{0.4}(\rho_a q_r)^{0.125}$  (Sun and Crook 1997)
- b) Reflectivity:  $Z_h = 43.1 + 17.5 \log(\rho_a q_r)$  (Sun and Crook 1997)



Fig. 1. Simulation domains. The green and black dots mark the best track of Typhoon Morakot (2009). The red dot and circle are the location and coverage of the CWB RCCG S-band Doppler radar.

ratio (more trustworthy)  $SNR_P = \frac{-42.7}{1}$   $SNR_Q = \frac{7.4}{1}$ Radar location is important for effectively observing the wind information

#### Assimilation cycles:

- 1) Updated model variables:  $V_r \rightarrow u, v, w \text{ and } q_r$
- $Z_h \rightarrow q_r$
- 2) Other variables also improve as the model integrates
- 3) RMS error and ensemble spread approach, although their trends are dominated by the large-scale dynamics that drives experiment NoDA



Fig. 2. The difference between the absolute values of the analysis and forecast errors in u, v, w (m s<sup>-1</sup>) and  $q_r$  (g kg<sup>-1</sup>) (left to right) at z=1 km when assimilating  $V_r$  at P and Q (top to bottom). The negative value (blue) represents improvement after assimilation while the positive value (red) represents degradation.



fig. 3. RMS error (black) and ensemble spread (red) during assimilation cycles for u, v, w (top row; left to right),  $q_r$ ,  $q_v$  and  $\theta$  (bottom row; left to right) averaged within radar coverage. The solid and dashed lines are of CTRL and NoDA, respectively.

### Deterministic forecast:

- 1) CTRL analysis at 18Z successfully retrieves the intensity and spatial pattern of the spiral rainbands, which are blurred in NoDA
- 2) Decaying capability to forecast the rainbands forming later than 18Z
  3) <u>QPF</u>: (Fig. 5)
  a) Both magnitude and pattern are improved for 3 hours, no matter in total or only torrential rain
  b) 3-hour improvement is fair due to the speed of the westerly

 $Z_h$  is modified to 0 dBZ if negative, where  $V_r$  is not available

# 3. Experimental design

- Simulated nature run and radar observations:
  - 1) Nature run: 00Z 8 Aug 00Z 9 Aug
    - a) Realistic rainfall compared with CWB observations
    - b) Rainbands comprising individual convective cells are well simulated
    - c) Windward slopes  $\rightarrow$  rain from the sea + terrain-induced convections Lee side  $\rightarrow$  drier downdrafts and less rainfall
    - d) Rainband evolution explains the analogous patterns of rainfall
  - 2) Observations: CWB RCCG radar
    - a) Realistic configurations  $\rightarrow$  7.5-min period, 9 sweeps, 230-km range
    - b) Simultaneous observations
    - c) Observation errors  $\rightarrow 1 \text{ m s}^{-1}$  for  $V_r$  and 2 dBZ for  $Z_h$
- Assimilation experiments
  - 1) Perfect model assumption
  - 2) Data assimilation is performed in the 3rd (finest) domain
  - 3) Assimilation run:
    - a) 12Z 8 Aug  $\rightarrow$  initialized and perturbed into a 30-member ensemble
    - b) 12-18Z 8 Aug  $\rightarrow$  model spin-up + assimilation cycles
    - c) 18Z 8 Aug 00Z 9 Aug  $\rightarrow$  deterministic forecast

Table 1. Assimilation strategies.



Fig. 4. Hourly maps of w (m s<sup>-1</sup>; colored) and the contours of  $q_r$ =1 g kg<sup>-1</sup> (gray lines) at z=1 km from 18Z to 21Z 8 Aug (top to bottom) for the nature run, CTRL and NoDA (left to right). P and Q mark the discussed single-point observations.

#### Horizontal localization test:

- 1) The CTRL forecast catches better intensity and position of the cell over terrain
- The forecast with doubled horizontal localization has a spurious cell over the sea



Fig. 5. RMS error, spatial correlation coefficient, ETS and BIAS (above 14.6 mm h<sup>-1</sup>; CWB's extremely torrential rain alert) of the hourly QPF within radar coverage (left to right). The solid and dashed lines are of CTRL and NoDA, respectively.



Fig. 6. Vertical cross sections of  $q_r$  (kg kg<sup>-1</sup>) at AA' (see Fig. 4) at 18:30 UTC 8 Aug. From left to right are the nature run, CTRL forecast and the forecast with doubled horizontal localization.

# 5. Summary and future prospects

Name	Assimilation duration	Cycle period	Superobservation- to-grid ratio	Model variables updated by <i>V<sub>r</sub></i>	Model variables updated by $Z_h$	Assimilation of zero Z <sub>h</sub>	1
CTRL	$\sim 2$ hours	7.5 min	1/4	$u, v, w, q_r$	$q_r$	Yes	optima

## **References and acknowledgement**

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#### Summary:

- 1) In OSSEs on Typhoon Morakot (2009), the WRF-LETKF radar data assimilation system improves the magnitude and pattern of the short-term QPF over complex terrain for 3 hours
- 2) Sensitivity tests are performed to optimize the assimilation strategies
- 3) Radar location is important for effective observations
- 4) Dynamics at large scales (beyond radar coverage) decides the trend
- Future prospects:
  - 1) Performance in probabilistic QPF by stochastic (ensemble) forecasts
  - 2) Tasks about real-case studies:
    - a) Higher grid resolution
    - c) Radar data quality control
- b) Reflectivity operator considering iced) Adaptive inflation and localization