

A Multi-scale Analysis on the Genesis of Hurricane Karl (2010) and Tropical Storm Matthew (2010).

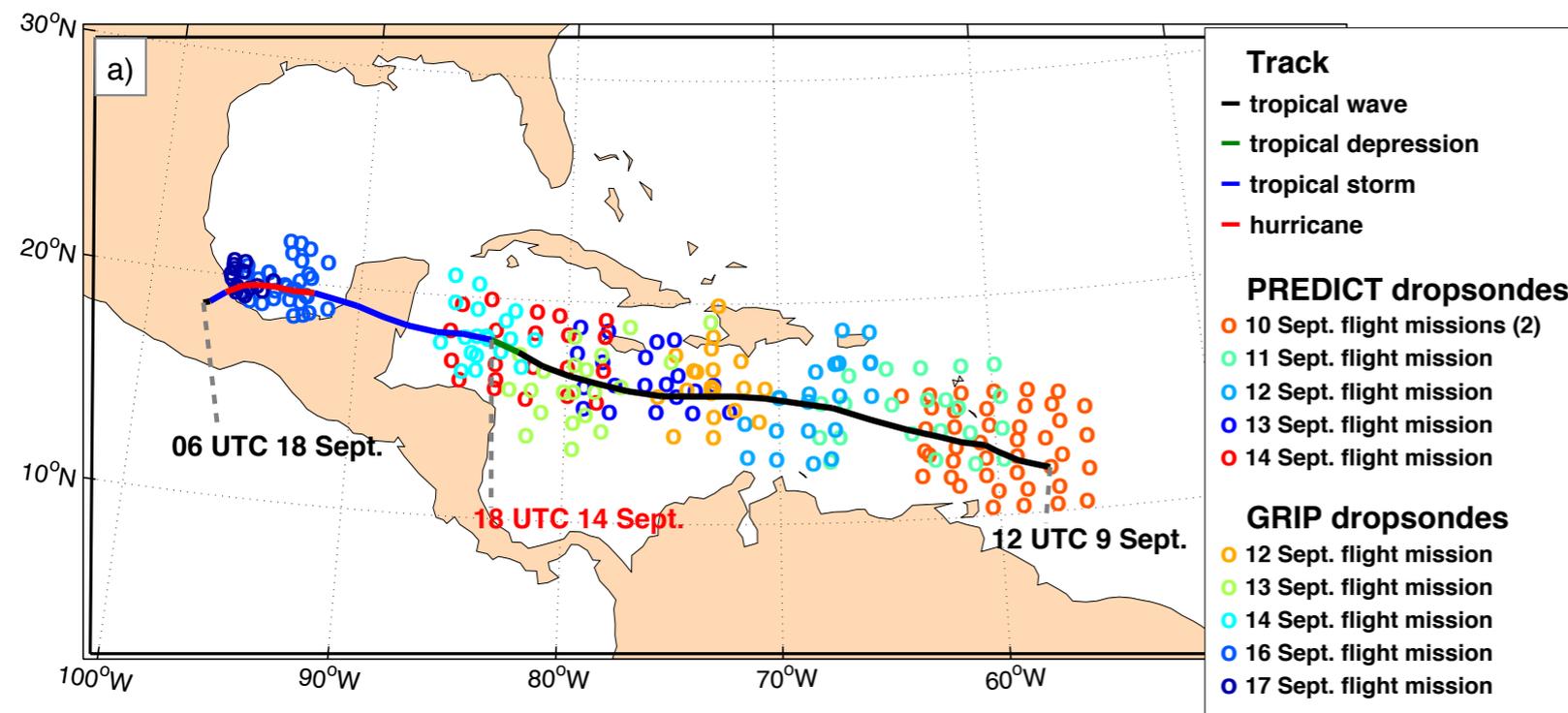
Ashford D'Arcy Reyes

Howard University Program for Atmospheric Sciences

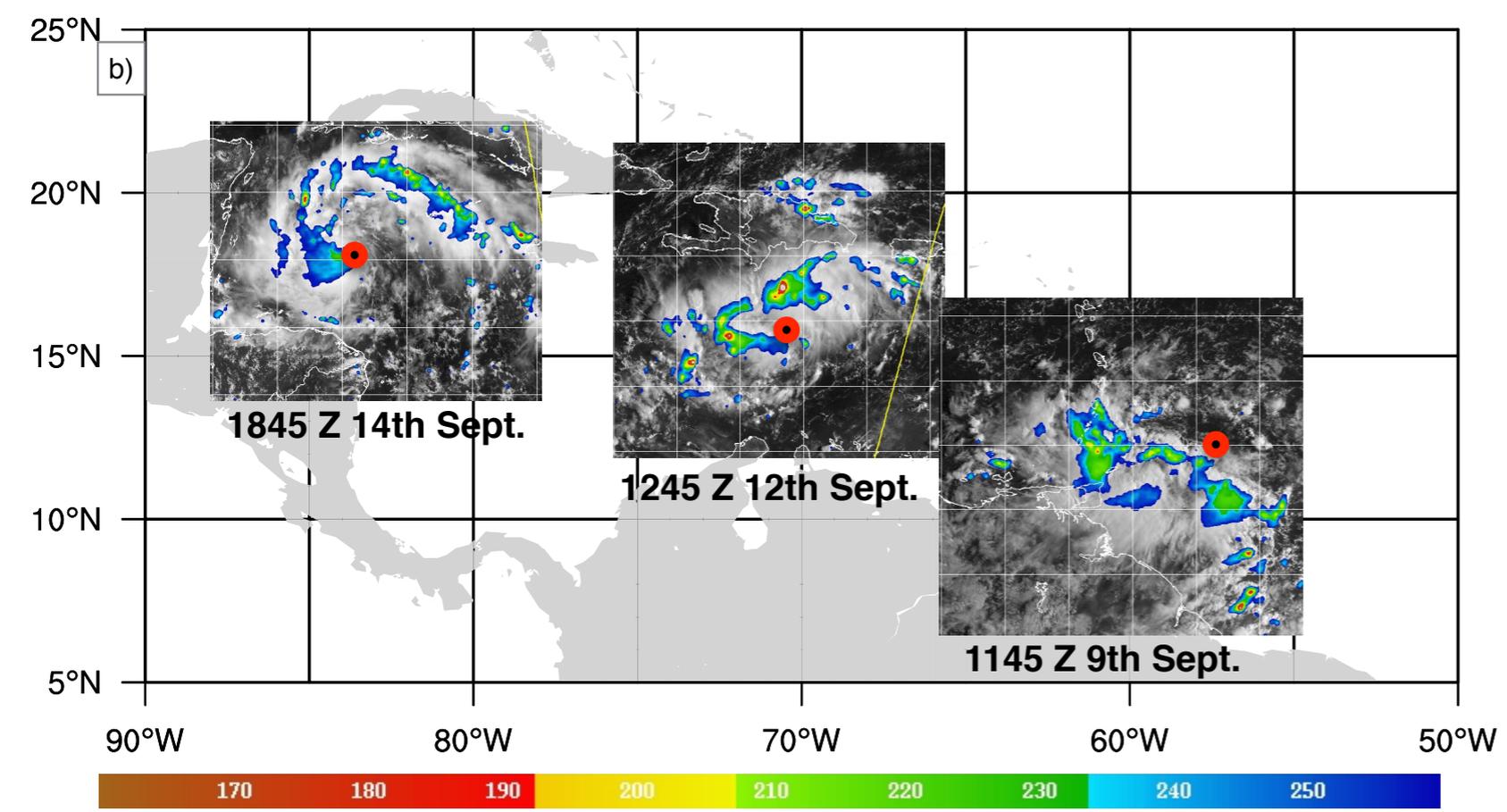
Spring 2015 Group Meeting

Friday 1st May 2015

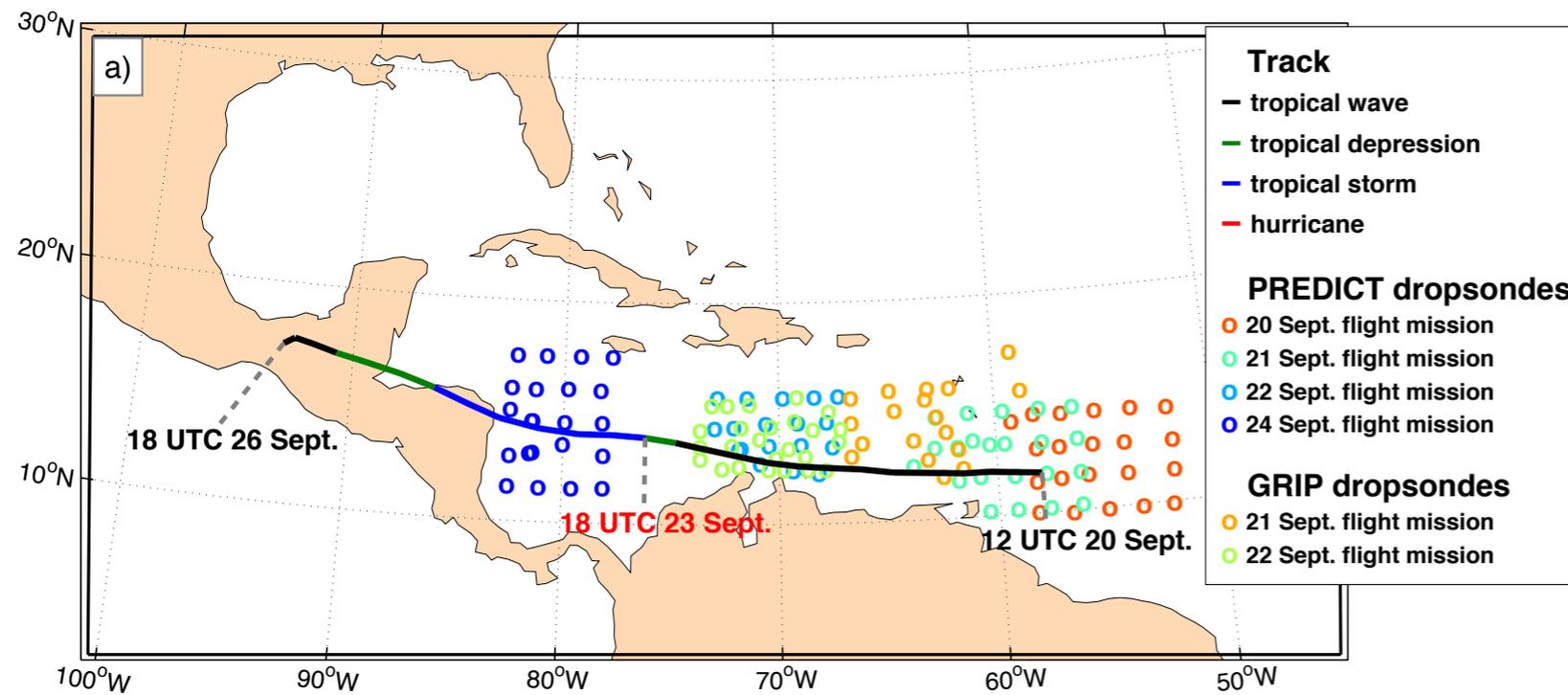
Hurricane Karl



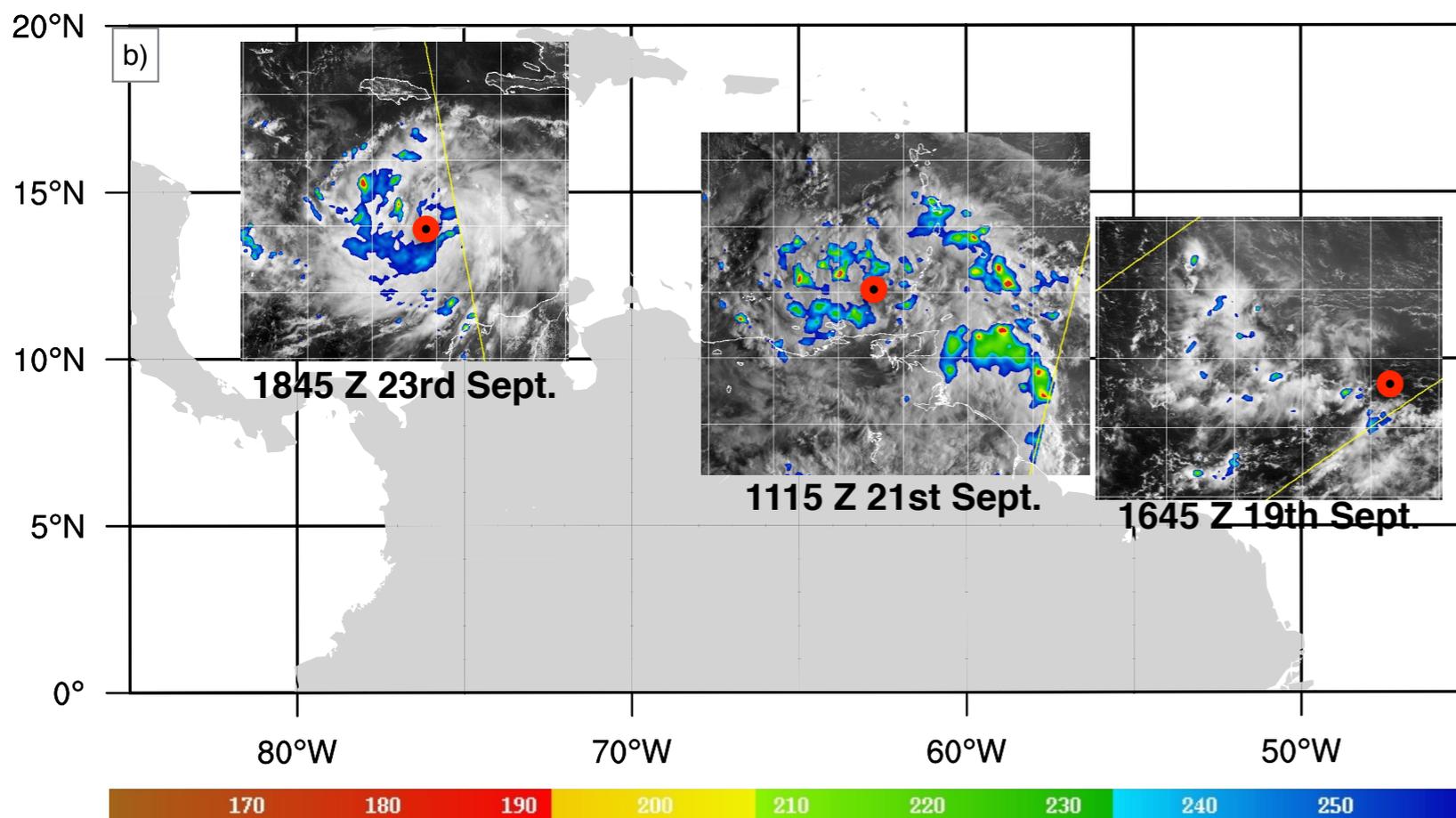
- Formed at 18.1 N, 83.6 W on September 14th at 18Z.
- Karl was a category 3 hurricane.
- Minimum central pressure of 956 hPa.
- Maximum sustained winds of 110 knots.



Tropical Storm Matthew

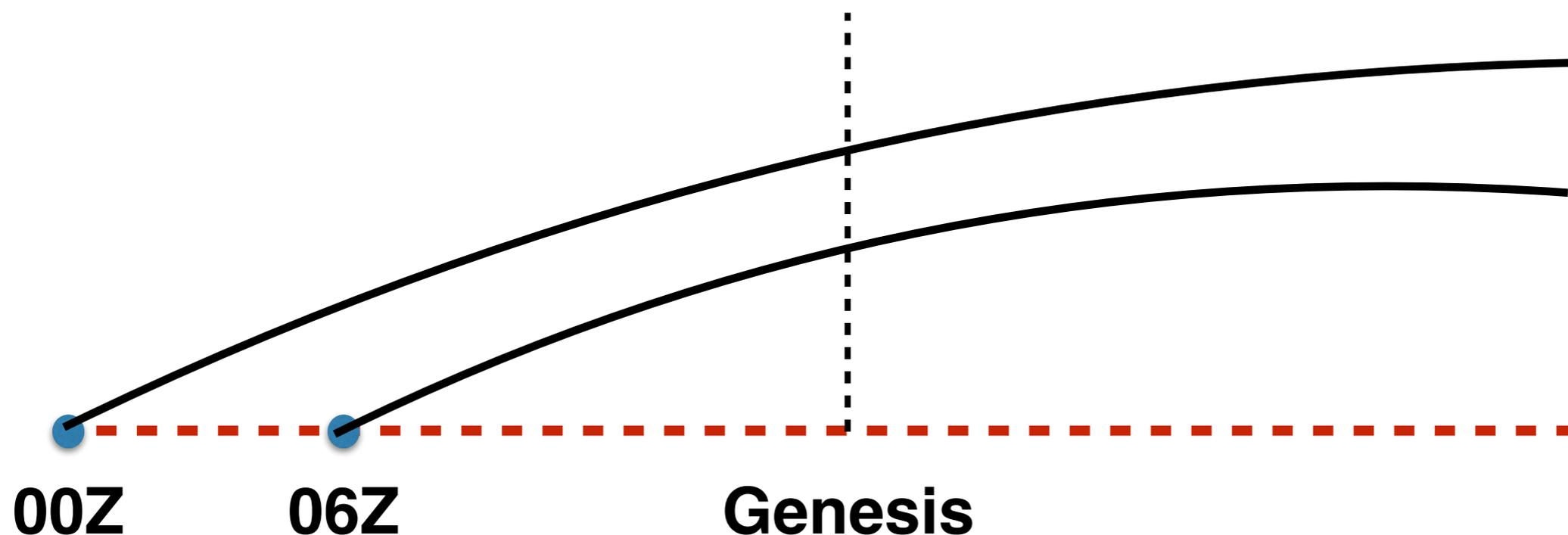


- Formed at 13.9 N, 76.2 W on September 23rd at 18Z.
- Matthew was a Tropical Storm (TS).
- Minimum central pressure of 998 hPa.
- Maximum sustained winds of 50 knots.



Experimental Design

- Experiments were initialized from GFS/GDAS analysis and cycled every 6 hours.
 - ▶ Karl initialized on Sept 8th 00Z and cycled through Sept 15th 00Z with deterministic forecasts until Sept 18th 06Z.
 - ▶ Mathew initialized on Sept 19th 00Z and cycled through Sept 24th 00Z with deterministic forecasts until Sept 26th 18Z.



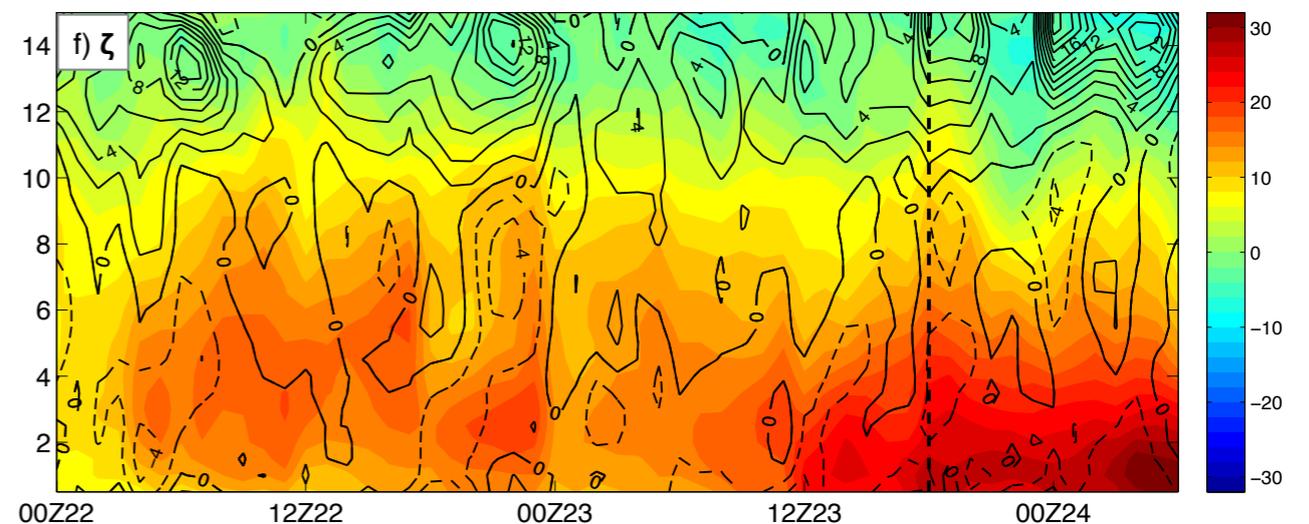
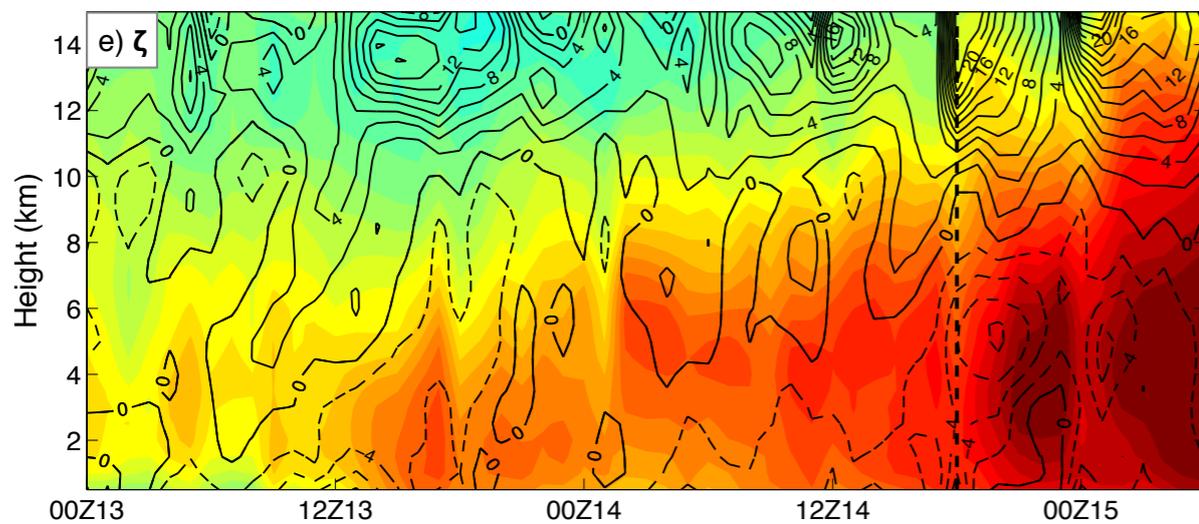
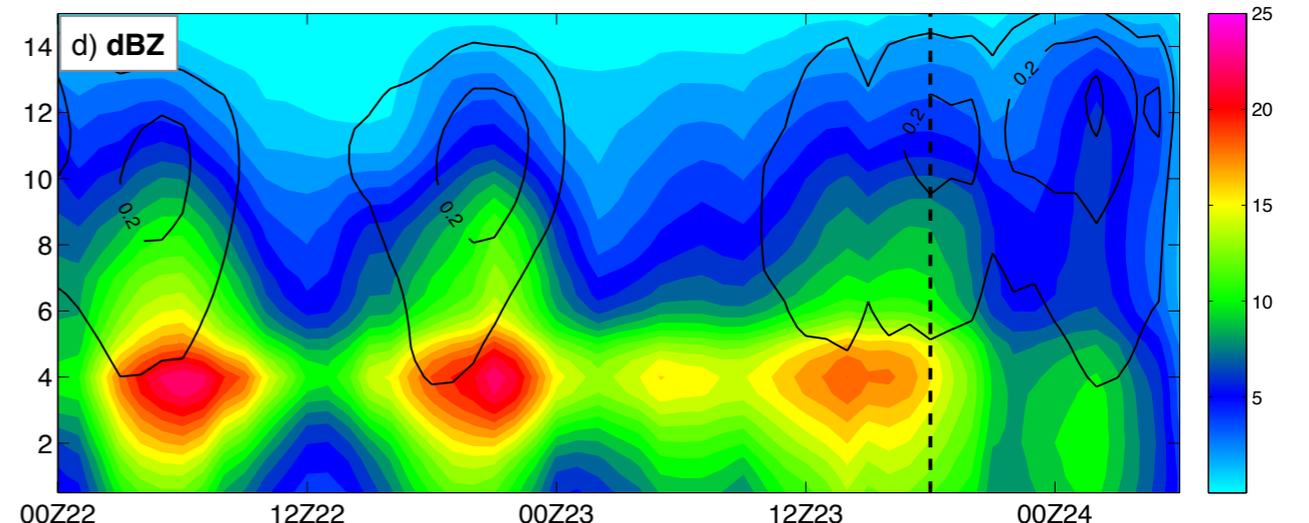
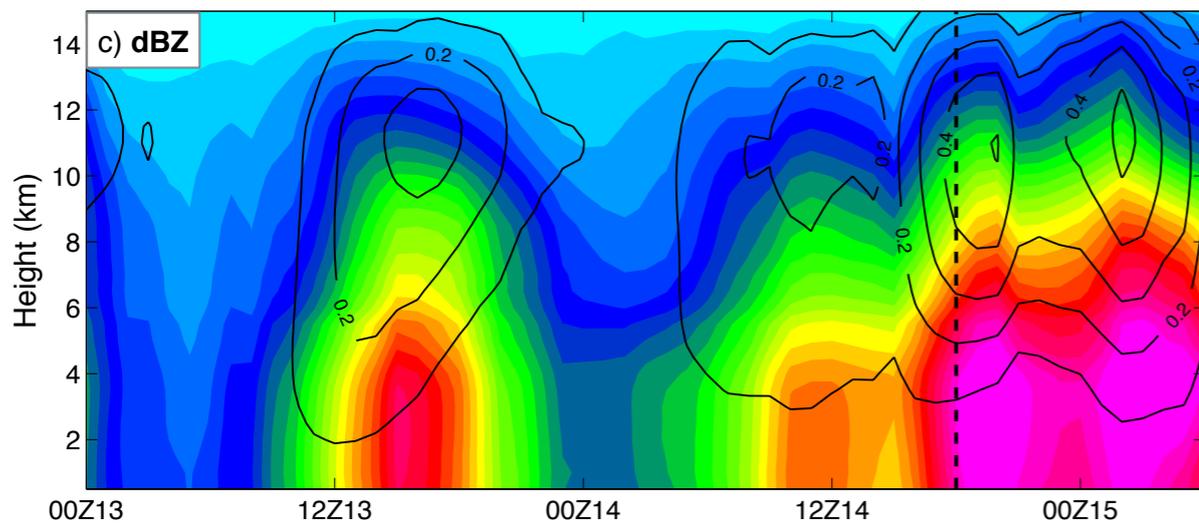
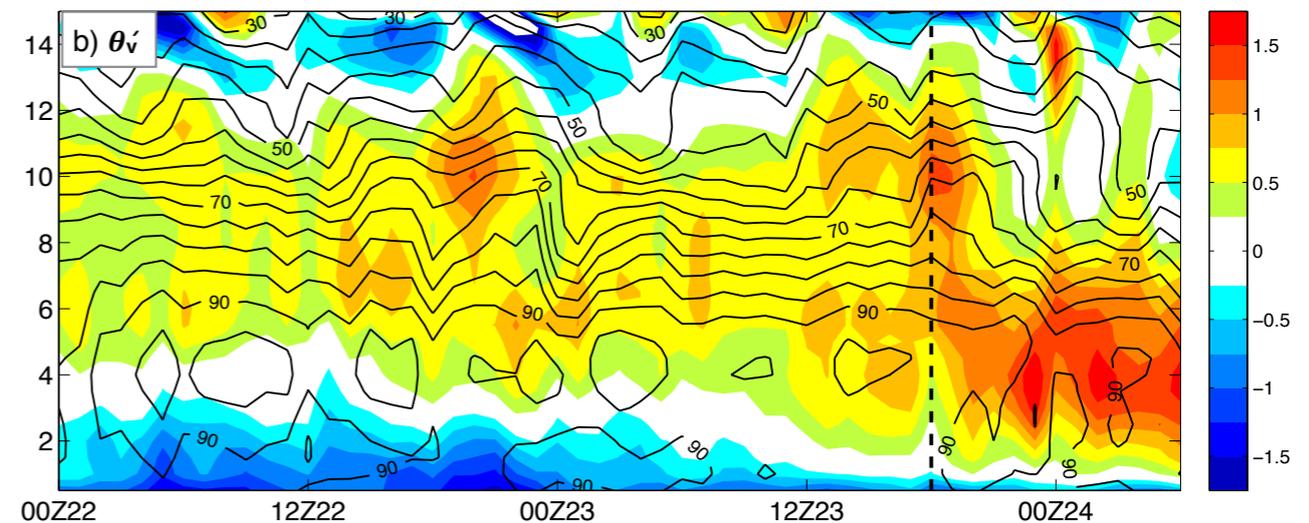
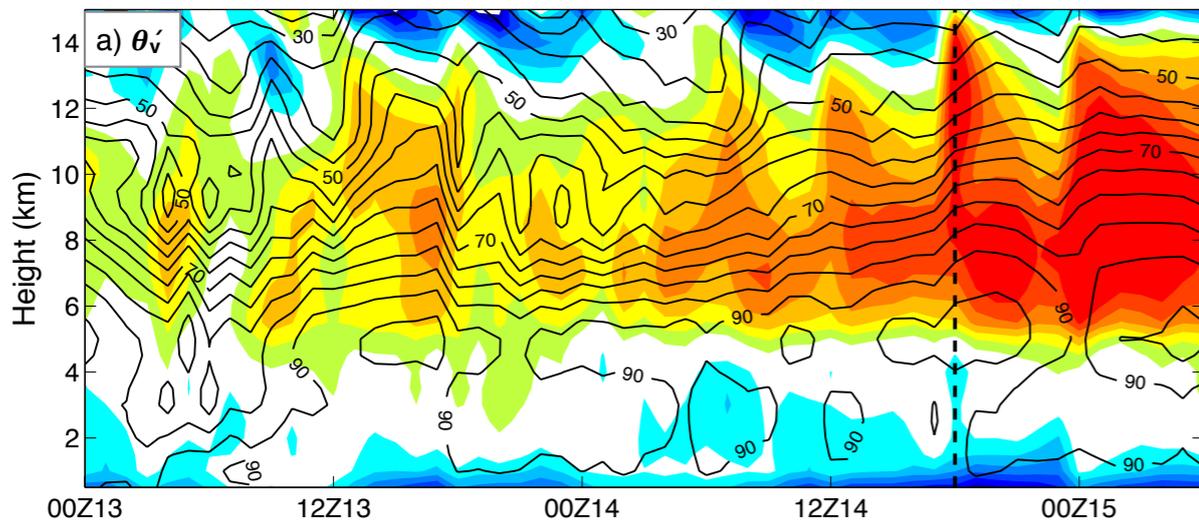
Model Setup

- Three domains with two-way nesting at 13.5-km, 4.5-km and 1.5-km grid spacing.
- 35 vertical levels together with a model top of 50 hPa.
- Explicit cumulus parameterization.
- Assimilation performed on 13.5-km domain.

Data Assimilation System

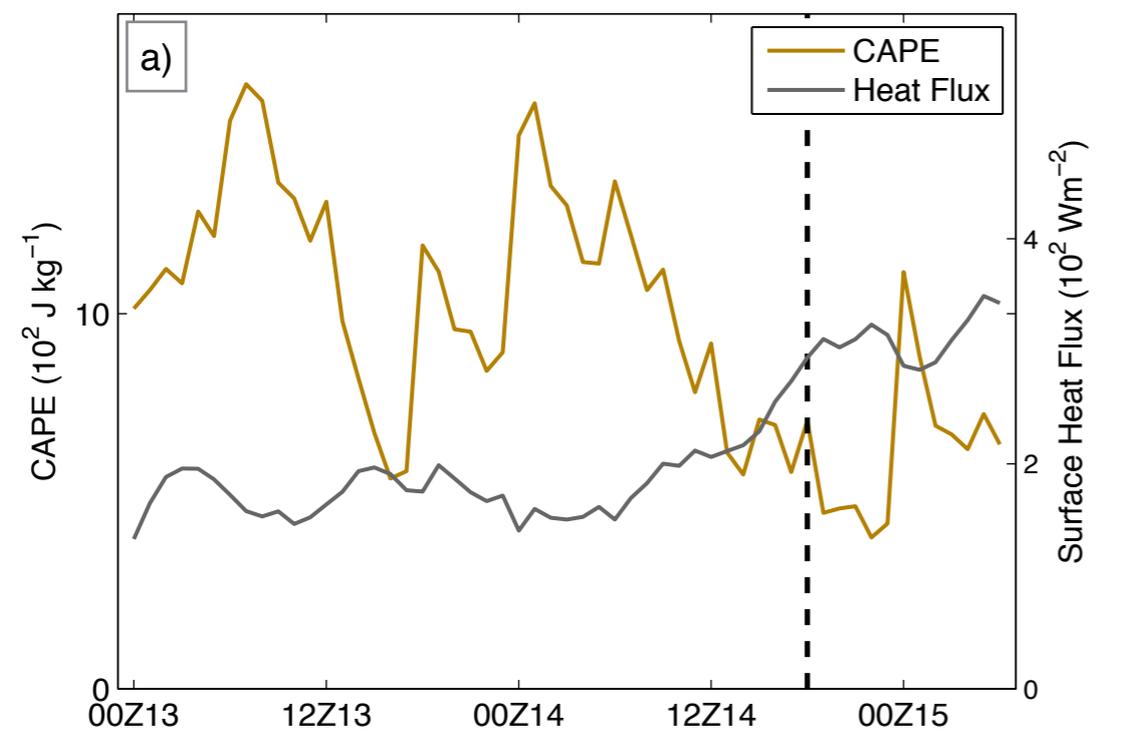
- 60 ensemble members
- Localization of 900 km in the horizontal & 15 levels in the vertical
- Relaxation coefficient of 0.8
- Two way coupling between EnKF and 4DVar
 - ▶ 4DVar uses ensemble mean first guess and ensemble perturbations
 - ▶ EnKF update the ensemble members
 - ▶ Hybrid 4DVar analysis replaces the EnKF analysis
- 80% of the increment comes from the ensemble perturbations during the hybrid minimization.

Inner-core (90 km) Vertical Profiles

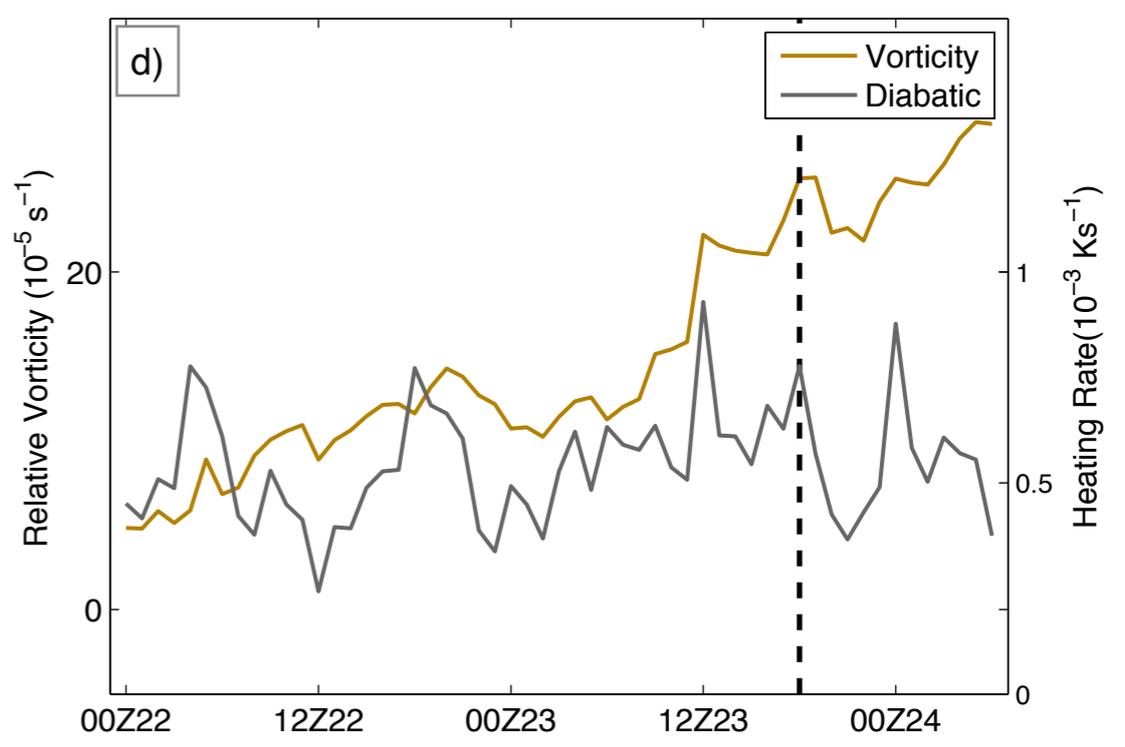
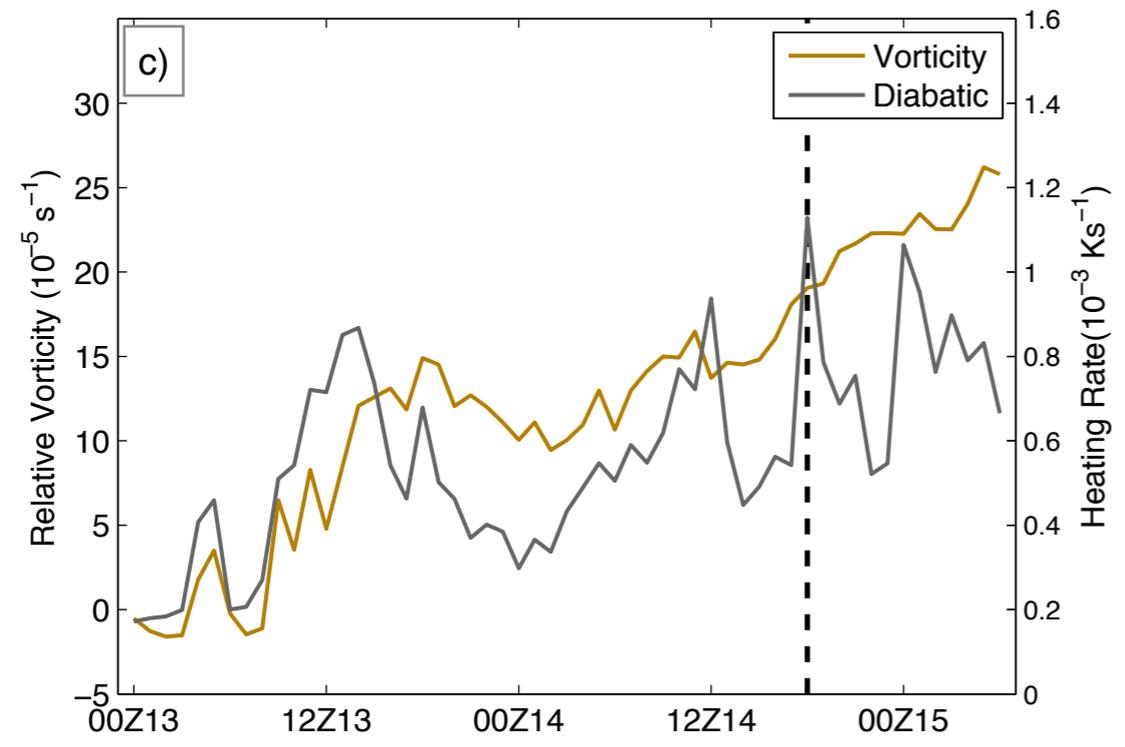
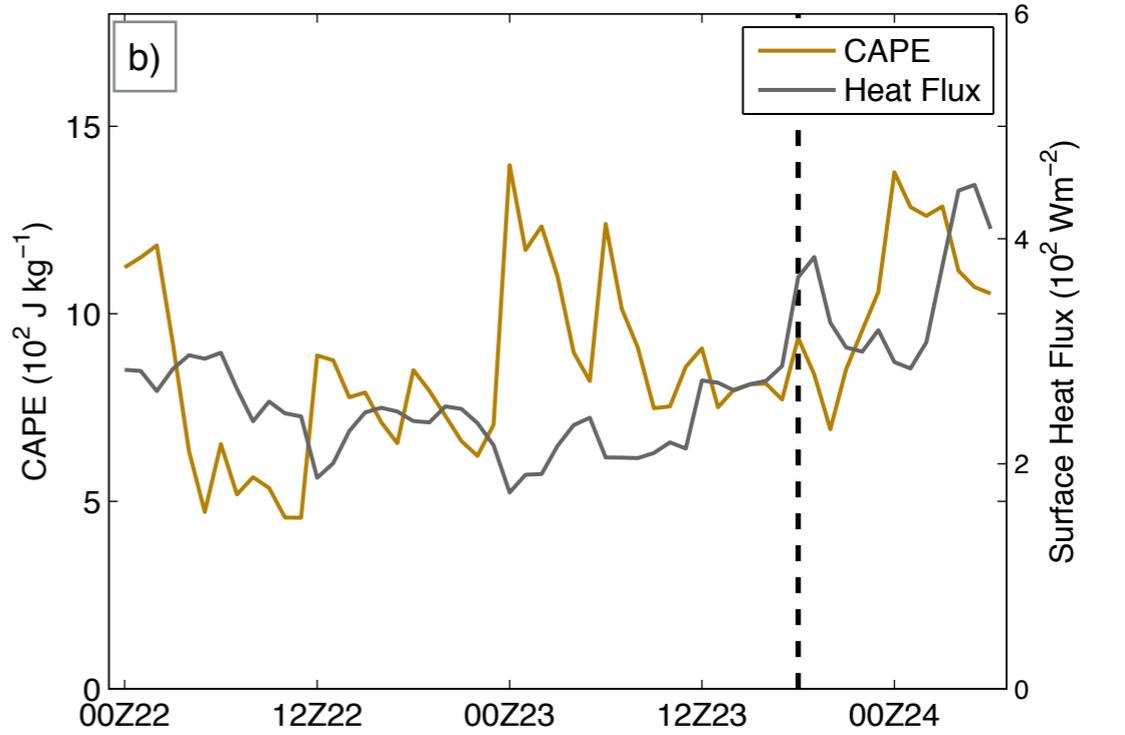
Karl
Matthew


Surface Heat Flux, CAPE, Vorticity and Diabatic Heating

Karl



Matthew



Vorticity Budget

Vorticity tendency equation:

$$\left. \frac{\partial \zeta}{\partial t} \right|_{SR} = -(V_h - C) \cdot \nabla_h (\zeta + f) - w \frac{\partial \zeta}{\partial z} - (\zeta + f) \nabla_h \cdot (V_h - C) - \left(\frac{\partial w}{\partial x} \frac{\partial v}{\partial z} - \frac{\partial w}{\partial y} \frac{\partial u}{\partial z} \right) + \frac{1}{\rho^2} \left(\frac{\partial \rho}{\partial x} \frac{\partial p}{\partial y} - \frac{\partial \rho}{\partial y} \frac{\partial p}{\partial x} \right) + \frac{\partial F_y}{\partial x} - \frac{\partial F_x}{\partial y}$$

Flux form of the vorticity tendency equation:

$$\left. \frac{\partial \zeta}{\partial t} \right|_{SR} = -\nabla_h \cdot \left\{ w \left[(\partial k \times V_h) / \partial z \right] \right\} - \nabla_h \cdot \left[(V_h - C)(\zeta + f) \right] + \frac{1}{\rho^2} \left(\frac{\partial \rho}{\partial x} \frac{\partial p}{\partial y} - \frac{\partial \rho}{\partial y} \frac{\partial p}{\partial x} \right) + \frac{\partial F_y}{\partial x} - \frac{\partial F_x}{\partial y}$$

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Vorticity Budget

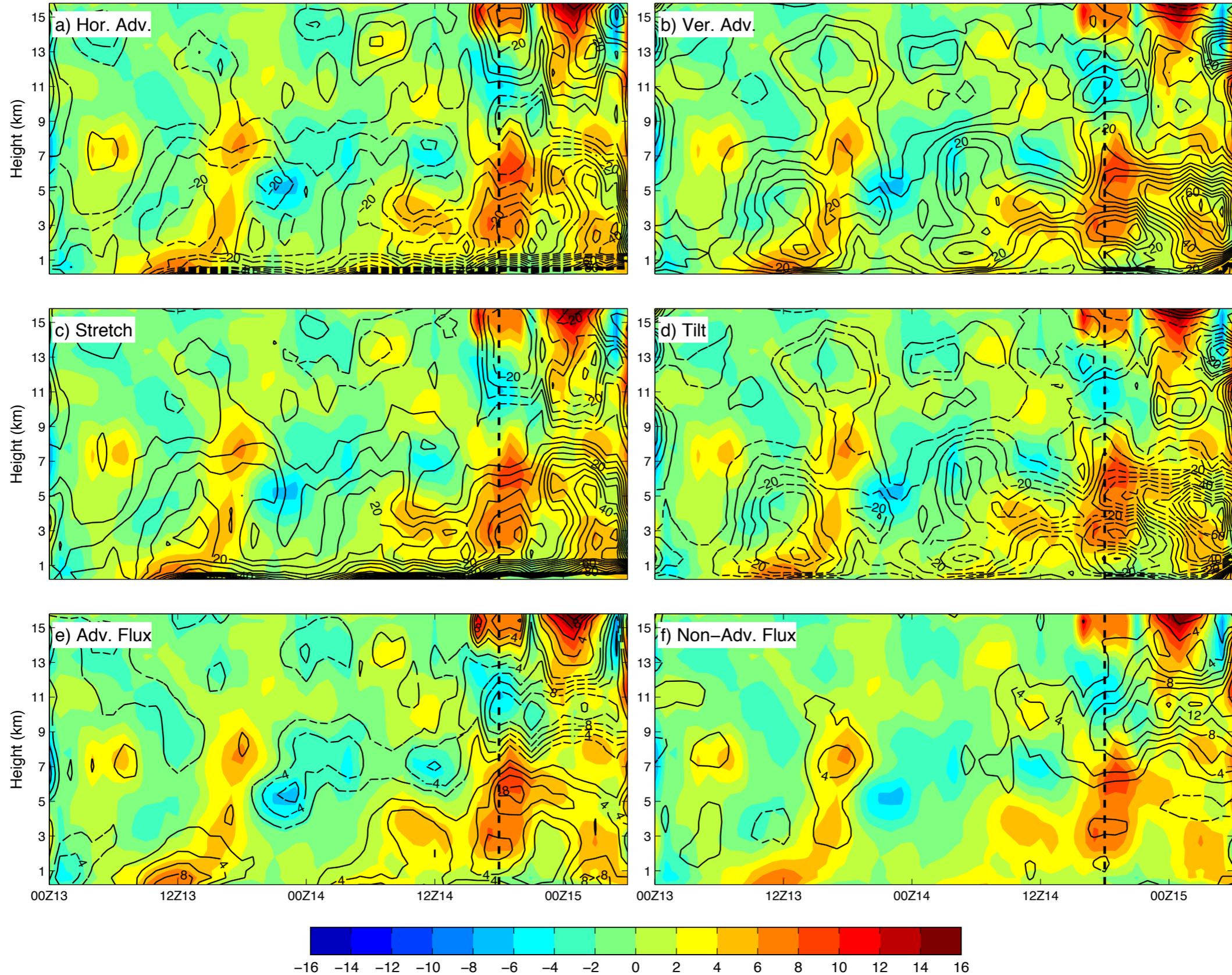
Vorticity tendency equation:

$$\left. \frac{\partial \zeta}{\partial t} \right|_{SR} = \underbrace{-(V_h - C) \cdot \nabla_h (\zeta + f)}_{\text{blue}} - \underbrace{w \frac{\partial \zeta}{\partial z}}_{\text{red}} - \underbrace{(\zeta + f) \nabla_h \cdot (V_h - C)}_{\text{blue}} - \underbrace{\left(\frac{\partial w}{\partial x} \frac{\partial v}{\partial z} - \frac{\partial w}{\partial y} \frac{\partial u}{\partial z} \right)}_{\text{red}} - \frac{1}{\rho^2} \left(\frac{\partial \rho}{\partial x} \frac{\partial p}{\partial y} - \frac{\partial \rho}{\partial y} \frac{\partial p}{\partial x} \right) + \frac{\partial F_y}{\partial x} - \frac{\partial F_x}{\partial y}$$

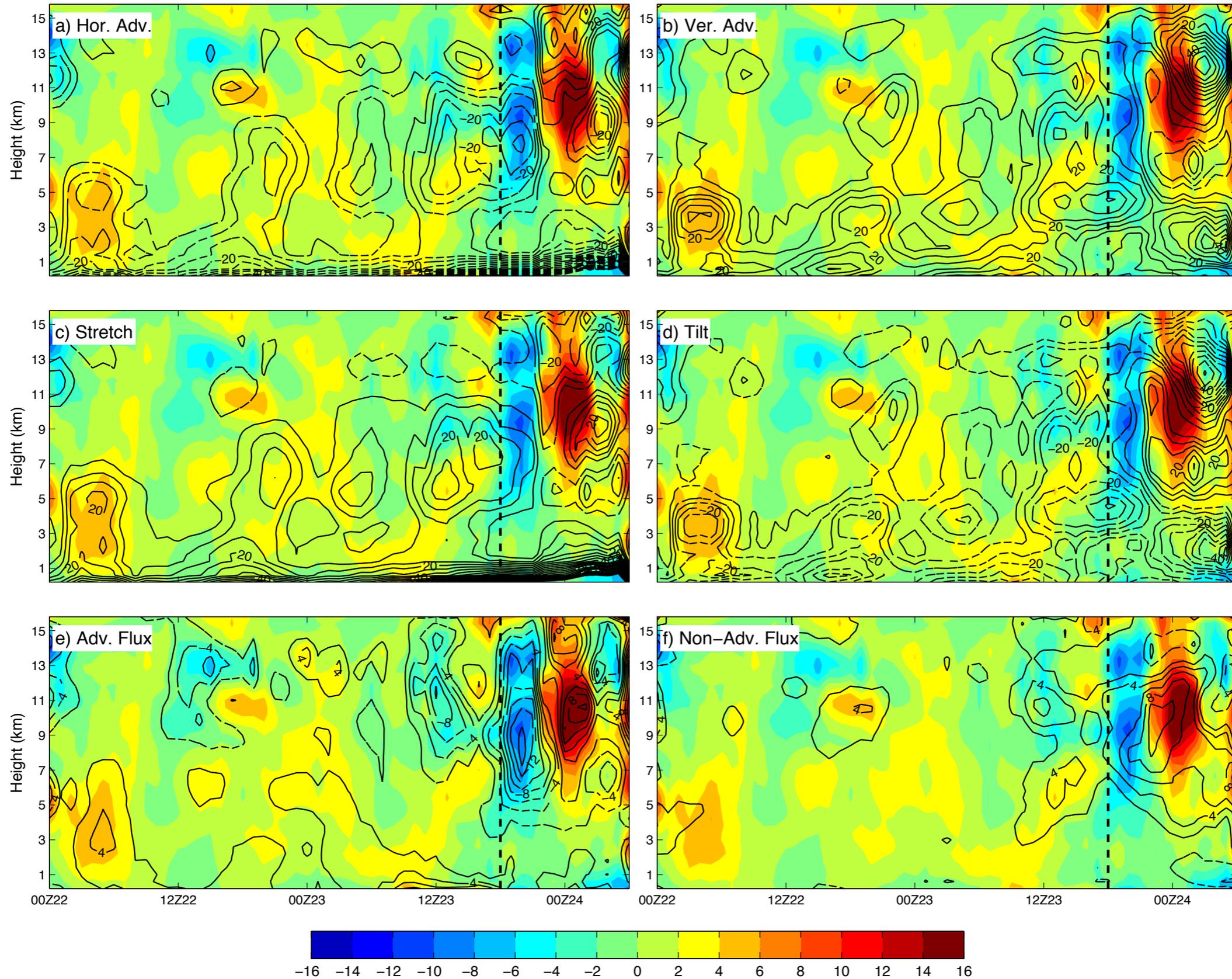
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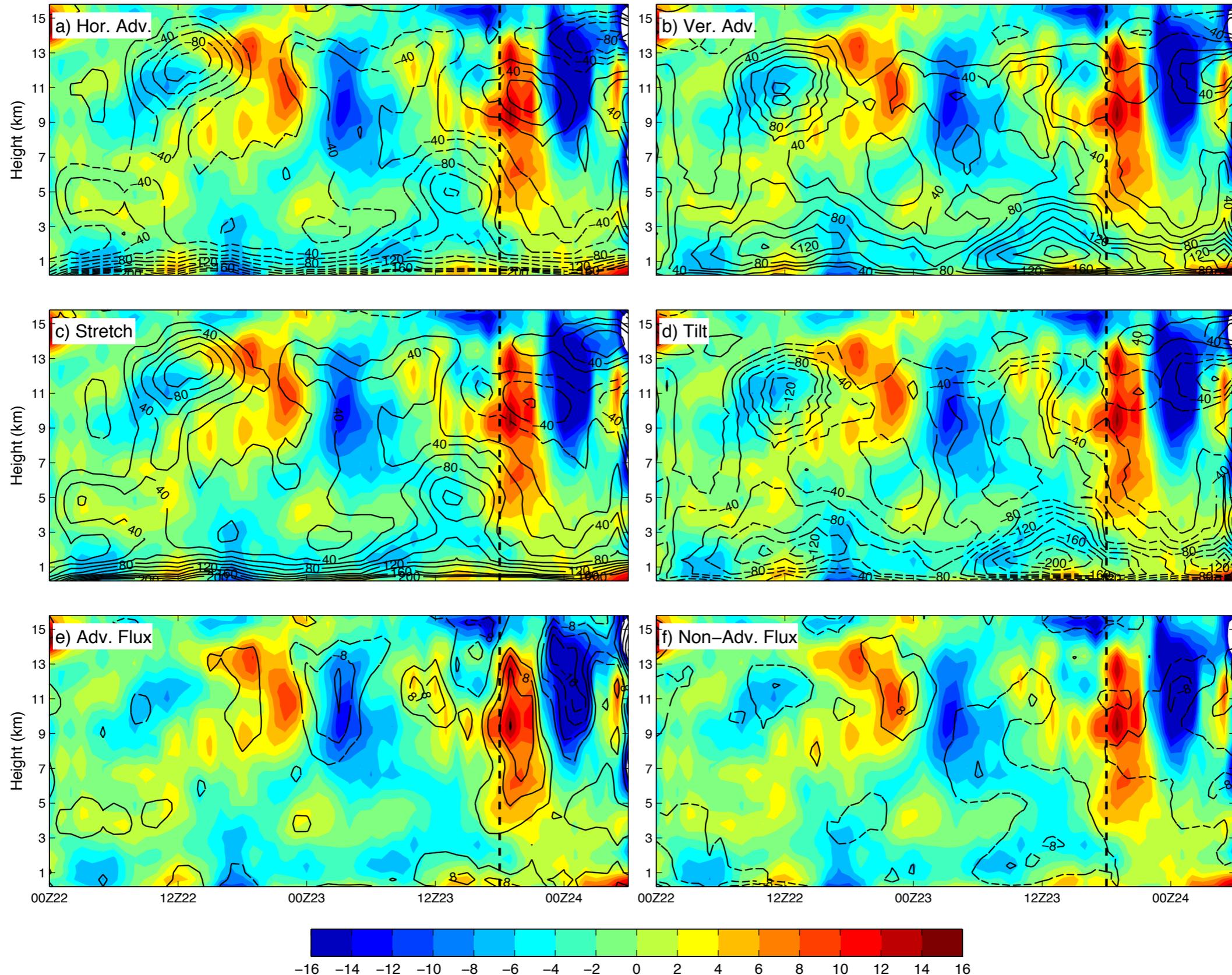
Hurricane Karl Vorticity Budget



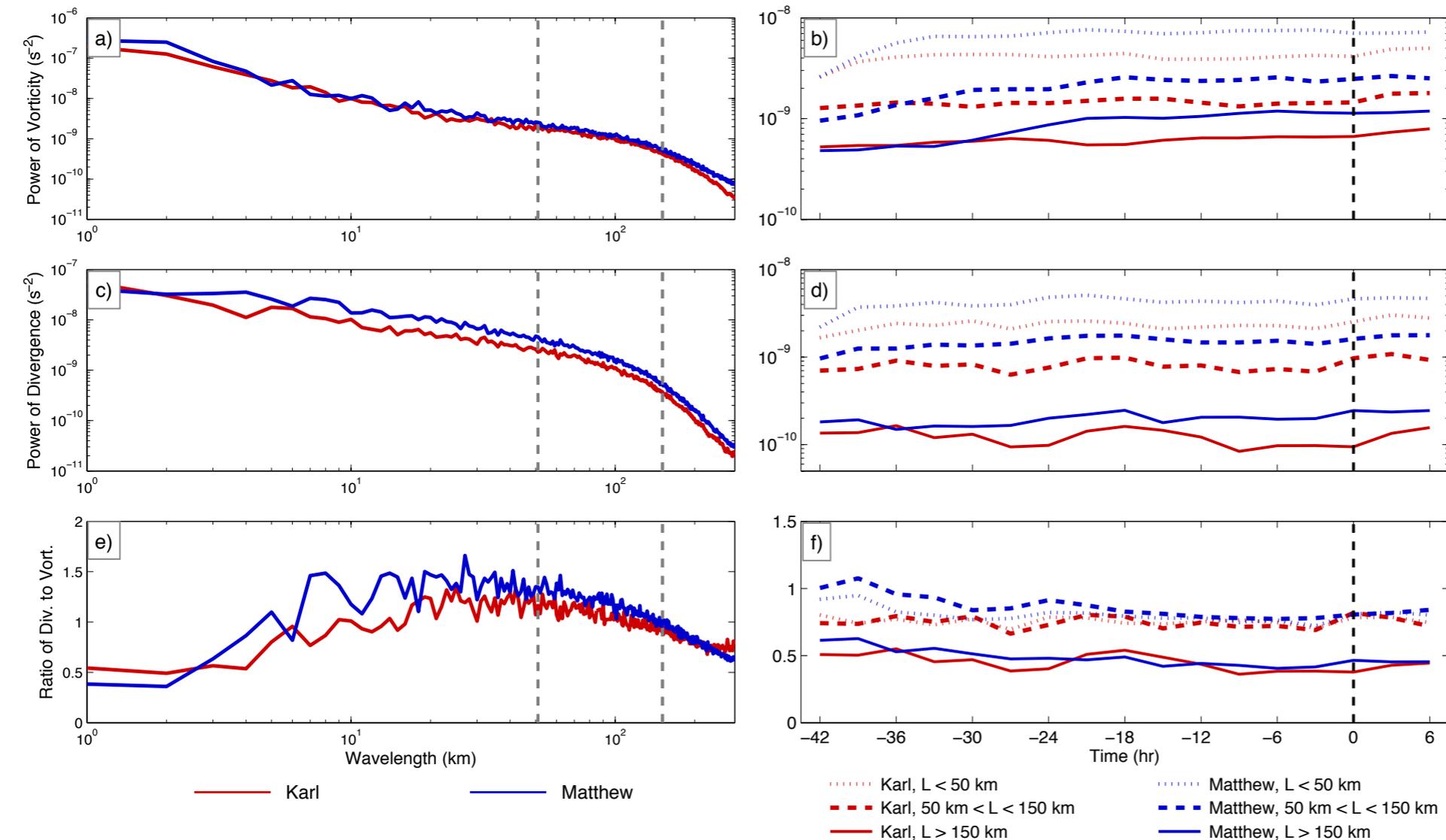
TS Matthew Vorticity Budget



TS Matthew Vorticity Budget (90 - 300 km)



Spectral Decomposition



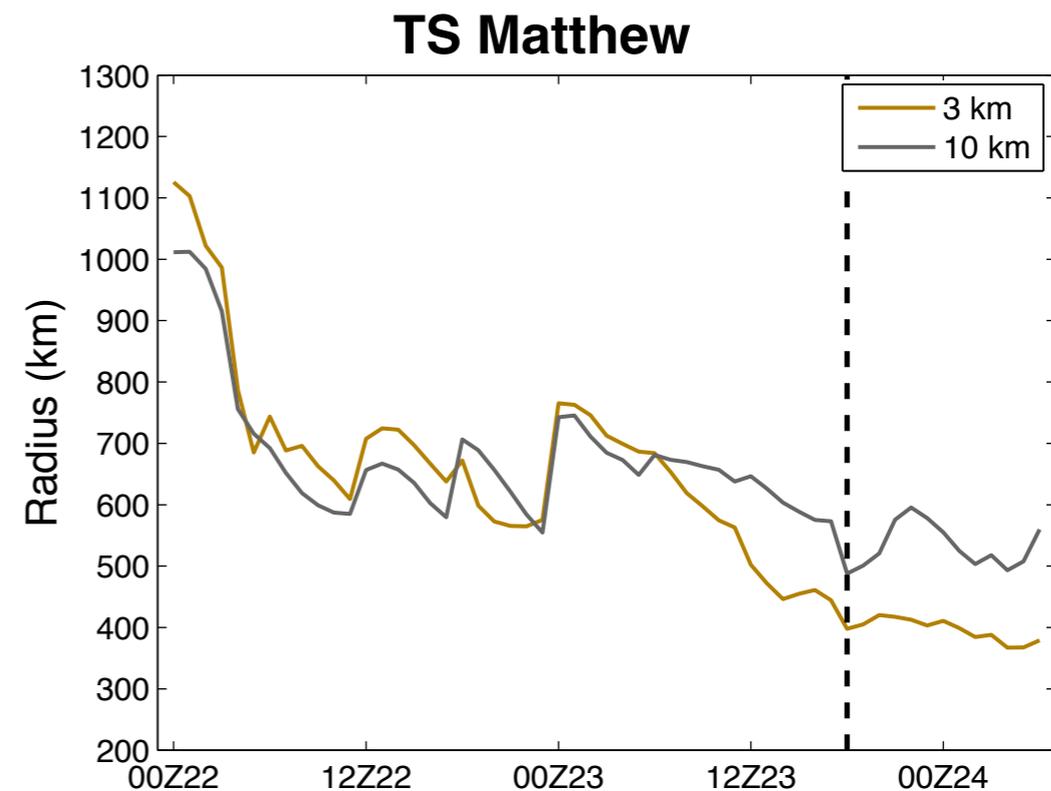
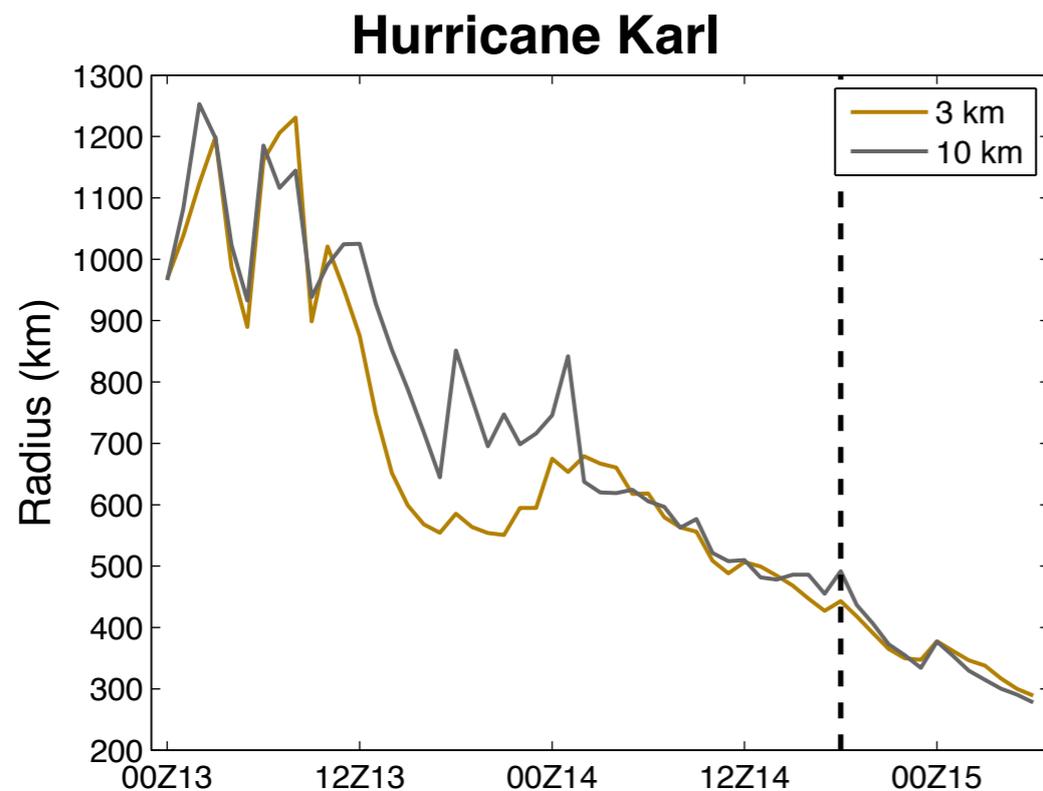
- Convective scale contribution is greater than the cluster and system scales.

- The square root of the ratio of Divergence to Vorticity shows the system scale is relatively balanced.

- Analysis are averaged over 12 hours centered on genesis.
- Convective scale : $L < 50$ km.
- Cluster scale : $50 < L < 150$ km.
- System scale : $L > 150$ km.

Rossby Radius of Deformation

$$L_R = \frac{NH}{(f + \zeta)^{\frac{1}{2}} (f + 2VR^{-1})^{\frac{1}{2}}} = \frac{NH}{I}$$



- Energy produced by convective heating within the lower to middle troposphere is confined to the system scale vortex.

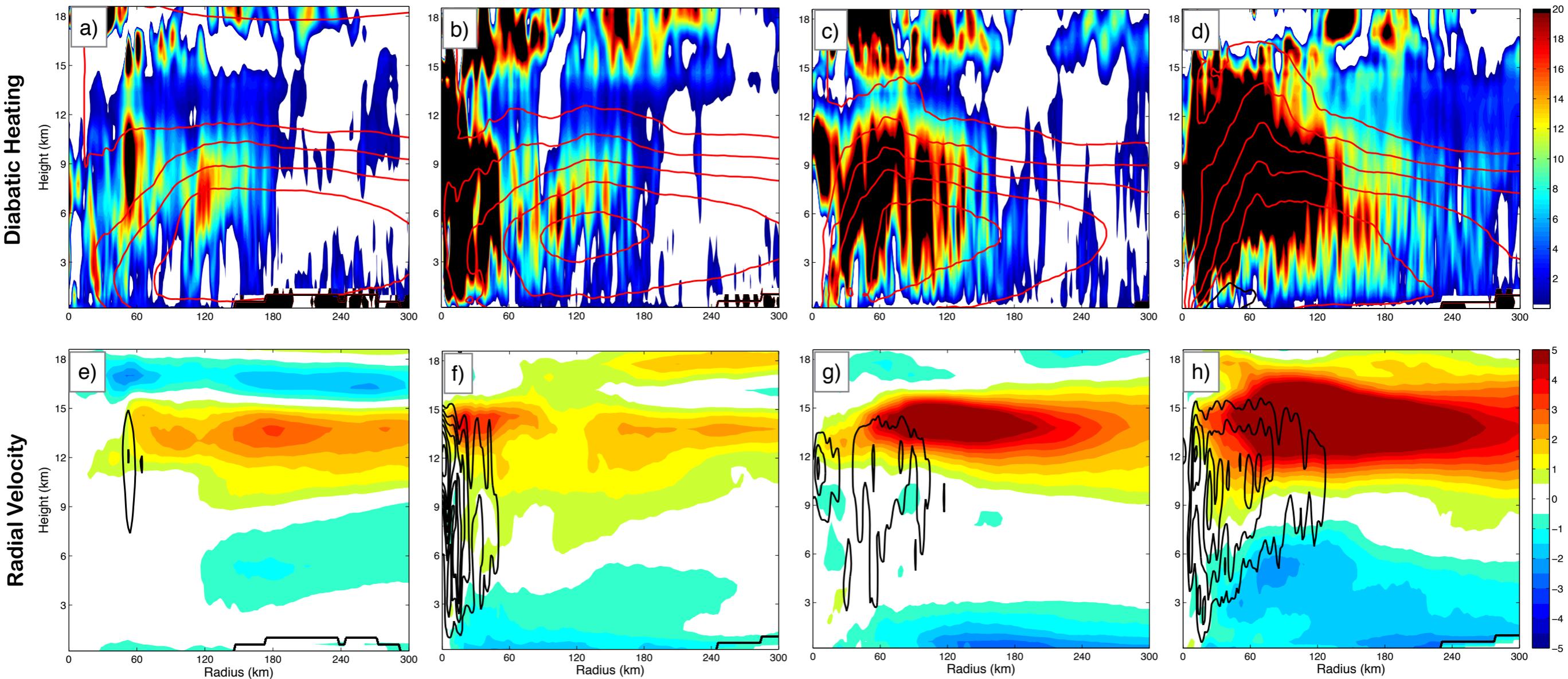
Hurricane Karl Secondary Transverse Circulation

24-hr forecast
00 UTC 14 Sept.

30-hr forecast
06 UTC 14 Sept.

36-hr forecast
12 UTC 14 Sept.

42-hr forecast
18 UTC 14 Sept.



- Heating and momentum forcing resulted in system scale convergent inflow in the lower troposphere.

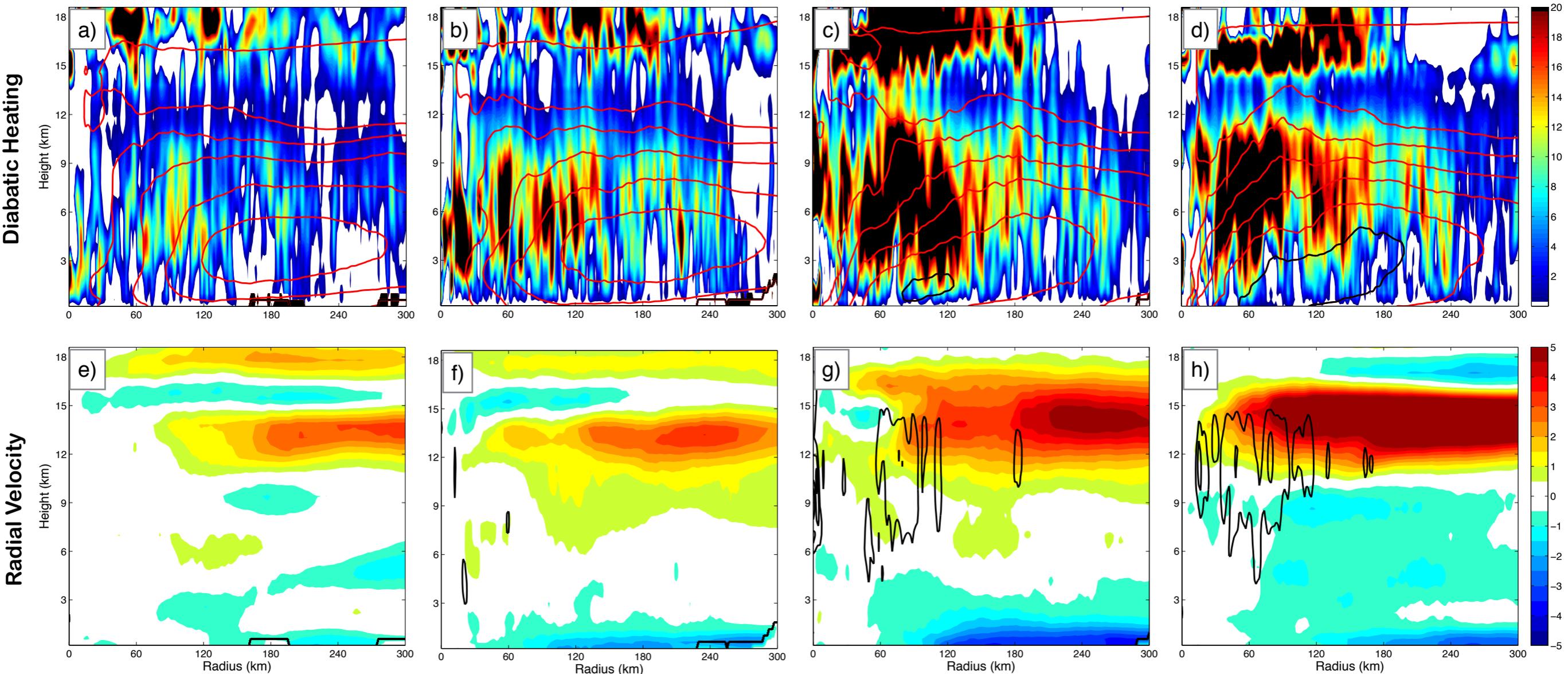
TS Matthew Secondary Transverse Circulation

24-hr forecast
00 UTC 23 Sept.

30-hr forecast
06 UTC 23 Sept.

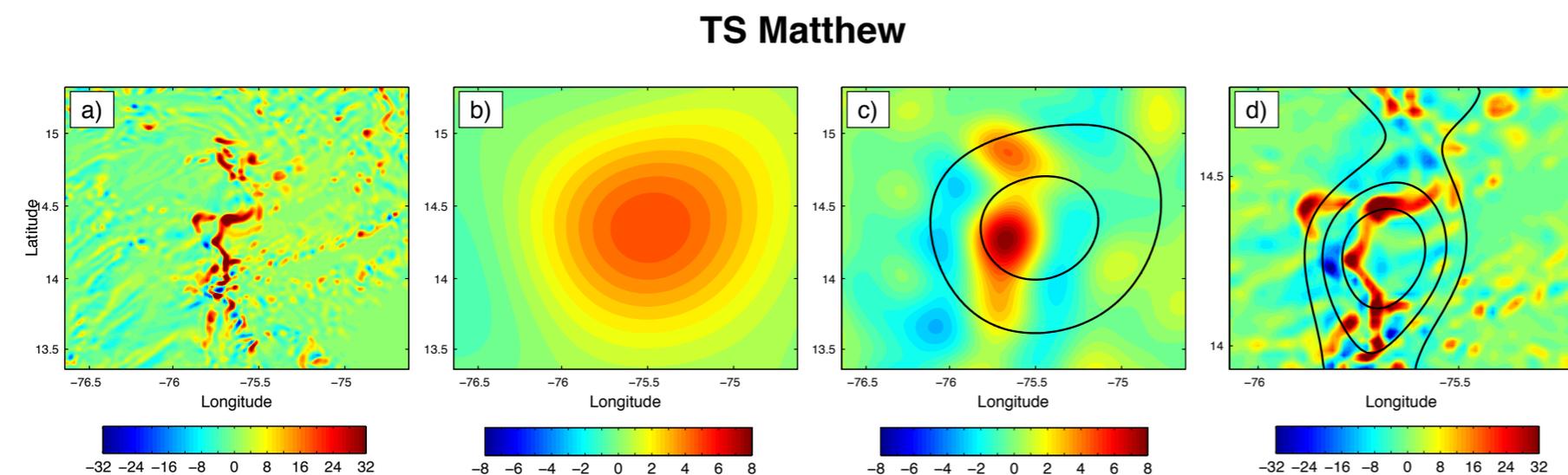
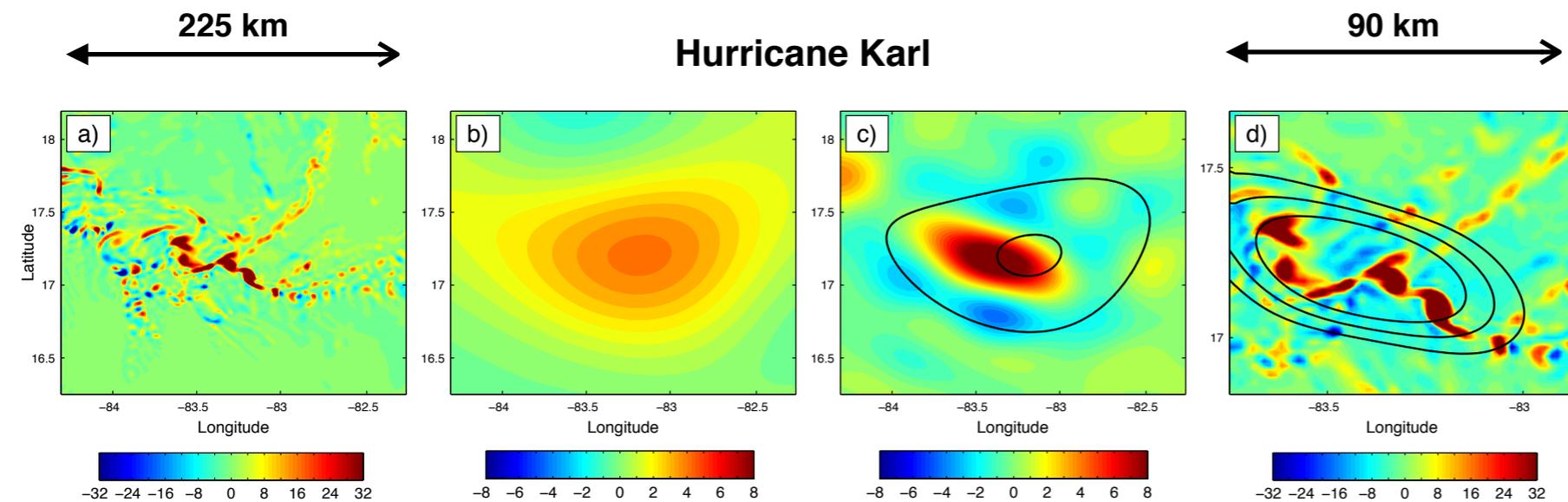
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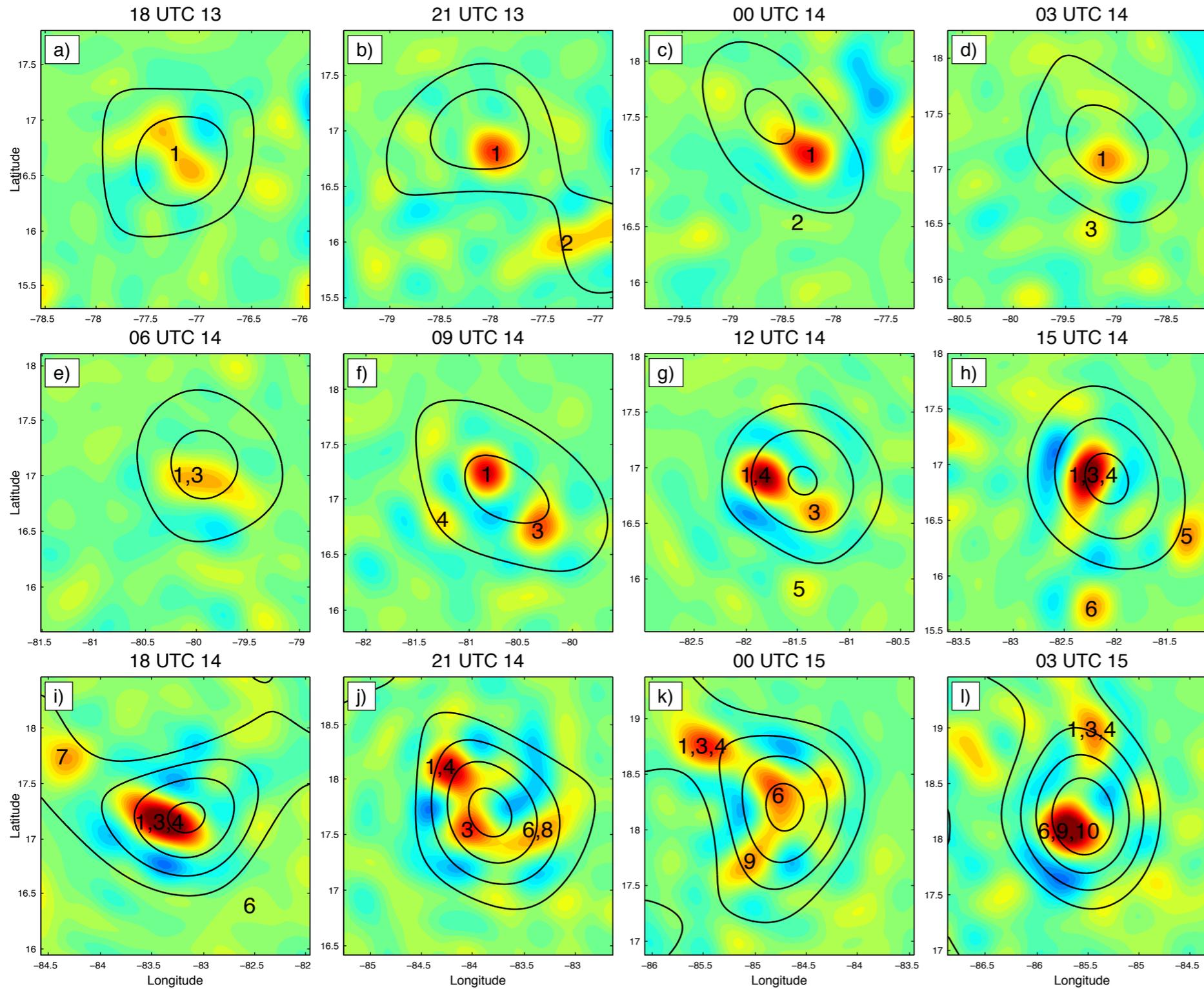
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Multi-Scale Features of a Developing Vortex at 200 meters



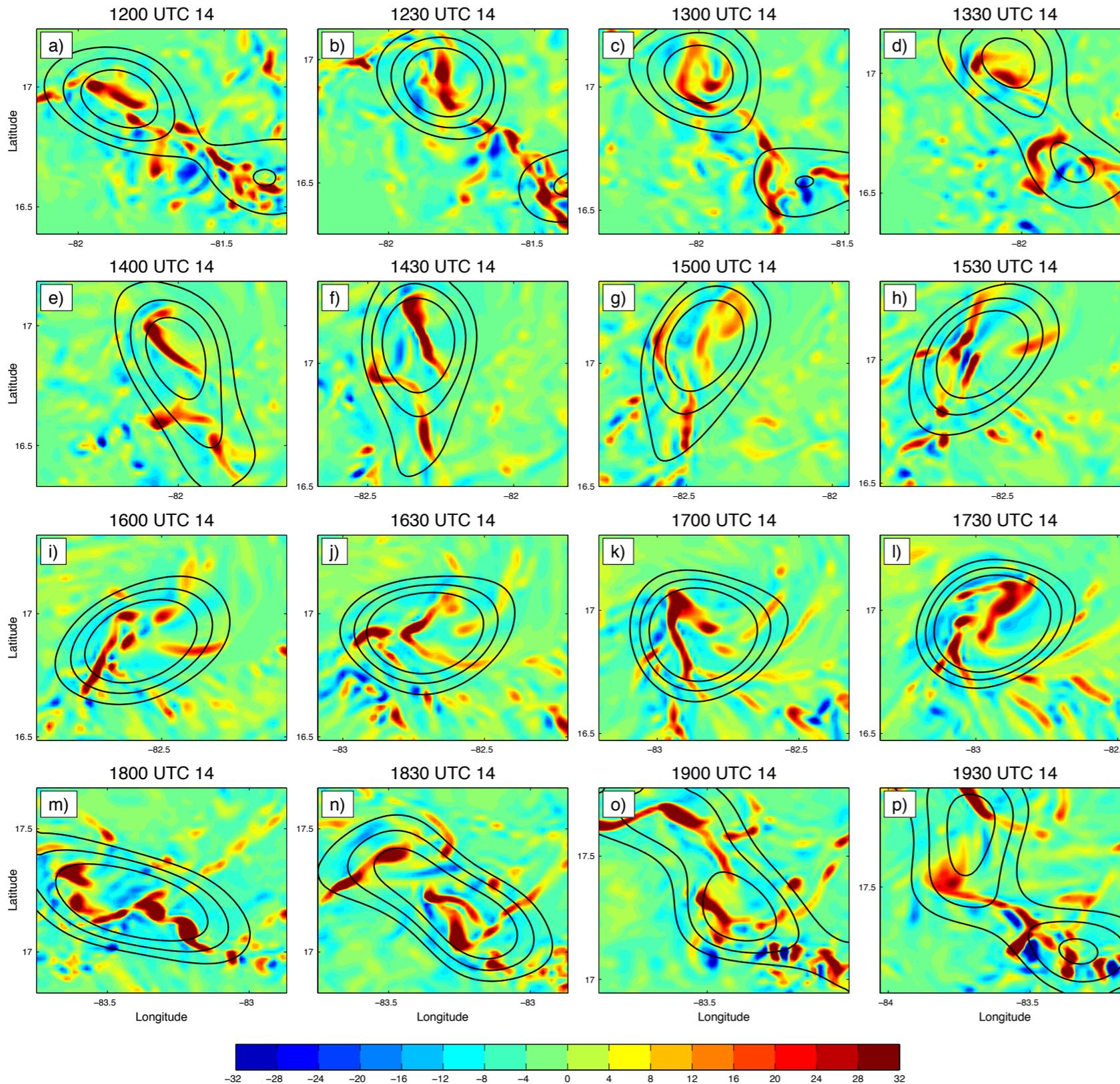
- Vorticity anomalies converge towards the center.
- Large number of CVAs within the cluster scale vorticity anomaly.
- Convective and cluster scale vortices are greater in magnitude than the system scale vortex.

Hurricane Karl Cluster Scale Vortex Evolution



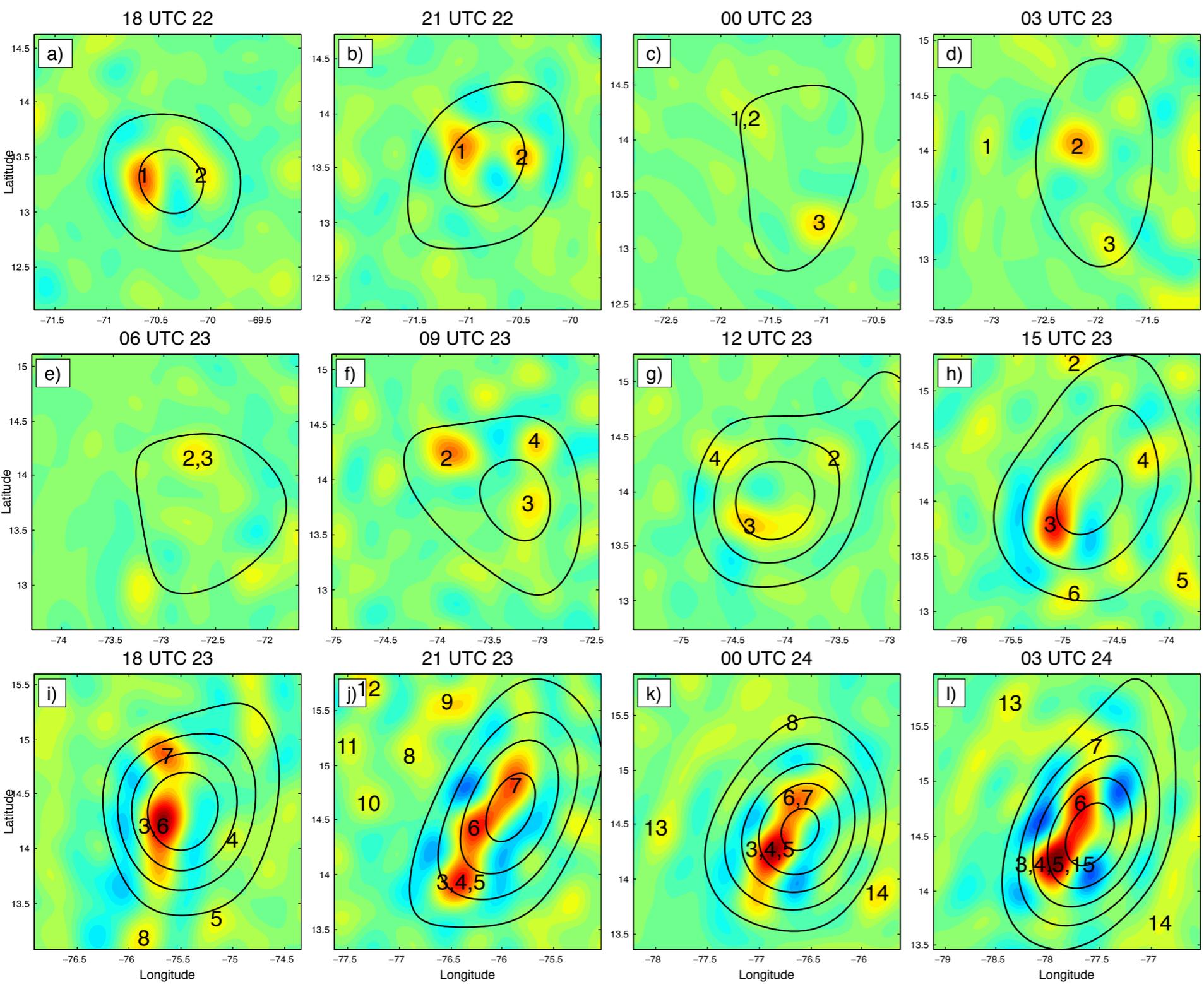
- Cluster scale vorticity mergers intensifies the system scale vortex.
- The merging of vortices 1,3 and 4 led to the intensification of the system scale vortex during genesis.

Hurricane Karl Convective Scale Vortex Evolution

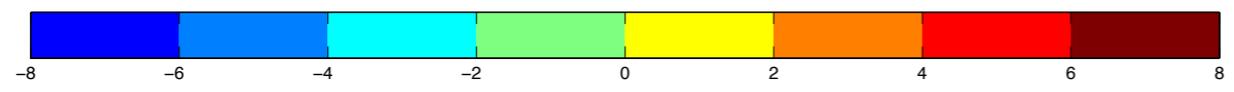


- The merger of CVAs within the cluster scale vortices 1,3 and 4 resulted in their mergers.
- Intensification of CVAs coincides with an intensification of the cluster scale vortices.

TS Matthew Cluster Scale Vortex Evolution



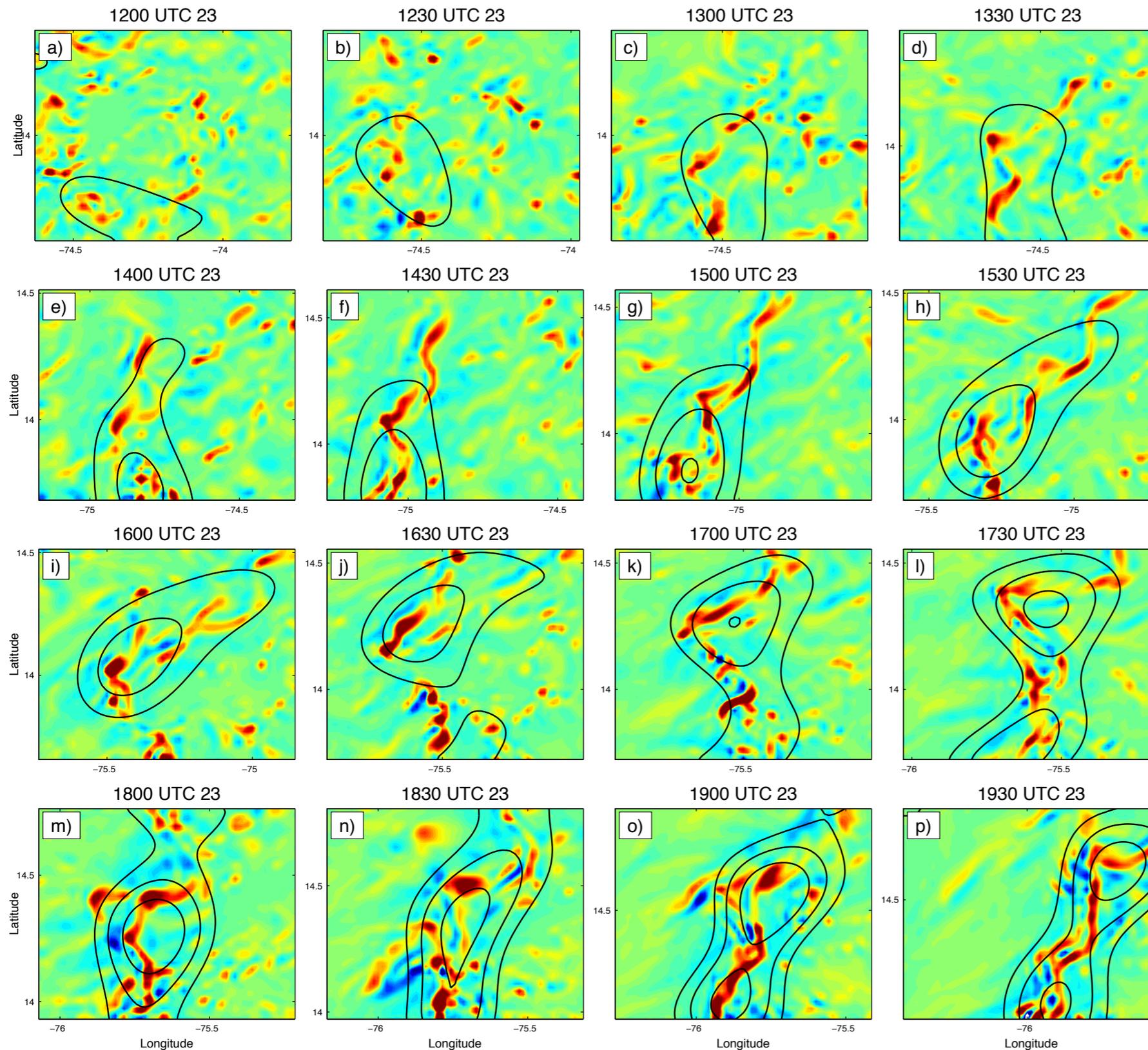
- Cluster scale vorticity mergers intensifies the system scale vortex.
- The merging of vortices 3,6 and 7 led to the intensification of the system scale vortex during genesis.



TS Matthew Convective Scale Vortex Evolution

The merger of CVAs within the cluster scale vortices 3,6 and 7 resulted in their mergers.

Intensification of CVAs coincides with an intensification of the cluster scale vortices.



Summary of Results

- TS Matthew shows important processes related to the top-down development theory.
- Vorticity budget shows that the low-level vortex was generated by the stretching term for both Karl and Matthew.
- Cluster-scale and Convective-scale vortex mergers are extremely important during the genesis process.
- CVAs are primarily responsible for the intensification of the low-level system scale vortex.
- Analysis imply that both systems genesis was as a result of the bottom-up process.

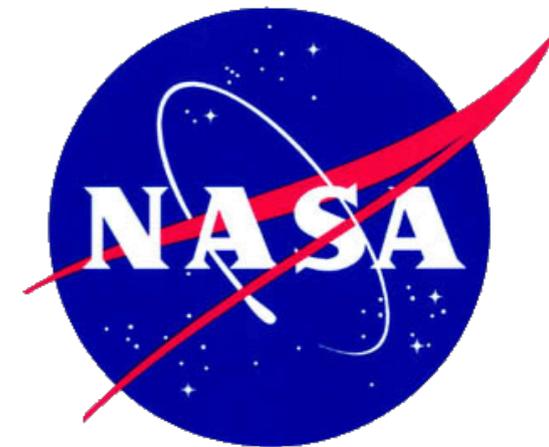
Future Directions

- Perform ensemble sensitivity analysis on Hurricane Karl and TS Matthew.
- Assimilate ground base radar observations to study the genesis of Hurricane Tomas (2010) and other mesoscale events affecting the Eastern Caribbean.
- Use a fully coupled atmosphere-ocean model like HWRF to study the multi-scale features of genesis.



Acknowledgments

PENNSSTATE



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Questions and Comments

